

The Composition and Crystal Structures of Pyrrhotite: A Common but Poorly Understood Mineral

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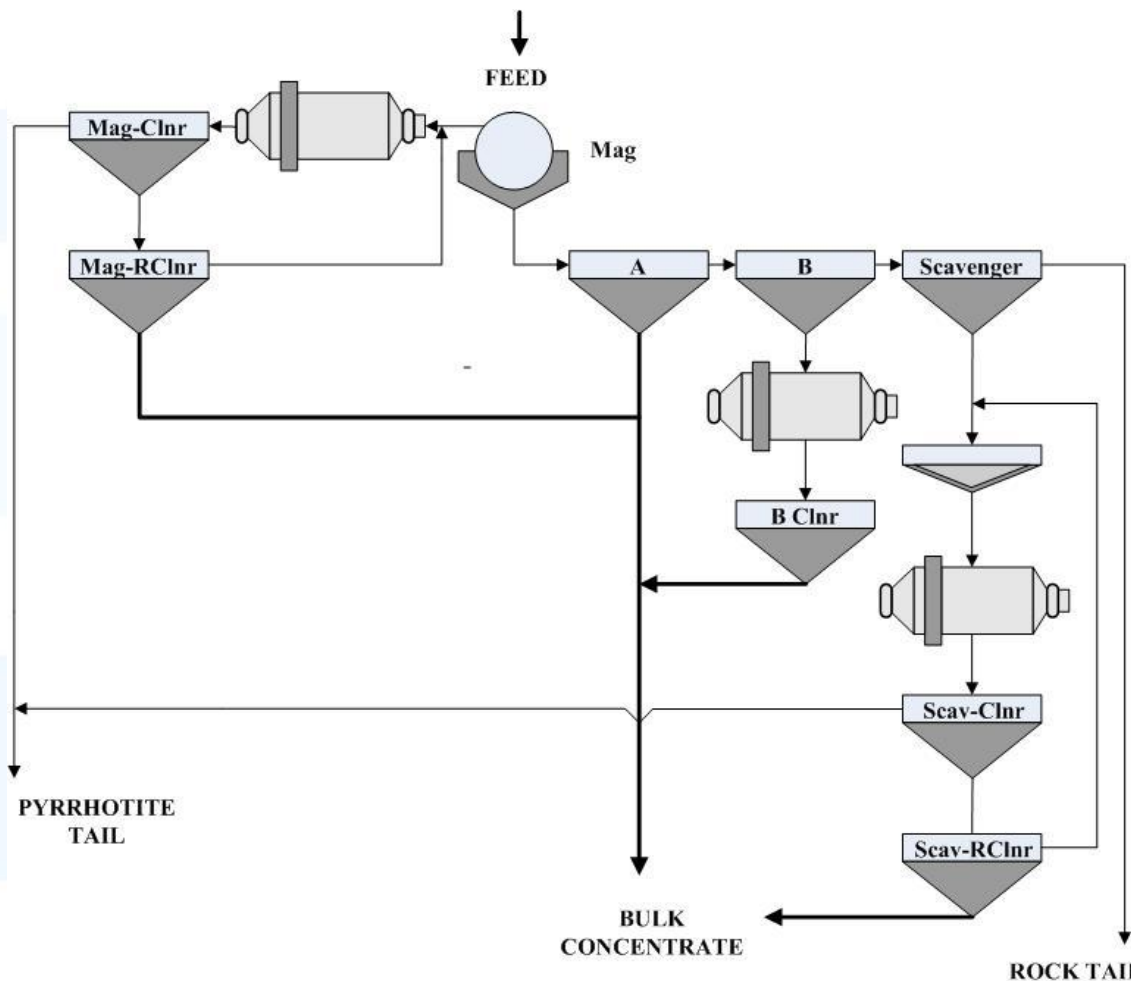
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Pyrrhotite Characterization

- Pyrrhotite (Fe_{1-x}S) is a common sulfide mineral in nickel and platinum group metal (PGM) deposits
- It is an unwanted phase in nickel processing because it dilutes the nickel concentrate and is expensive to remove
- It is recovered in PGM processing because many platinum group minerals are associated with it. It also contributes to matte-fall in UG-2 matte smelting
- Pyrrhotite occurs in several forms with different symmetry and usually slightly different compositions
- In Nature two major forms exist:
 - Magnetic Fe_7S_8 and can in principle be separated from.....
 - Non-magnetic Fe_9S_{10} and $\text{Fe}_{11}\text{S}_{12}$
- We are investigating the flotation behaviour of the two pyrrhotite types – hence the need to characterise and quantify them



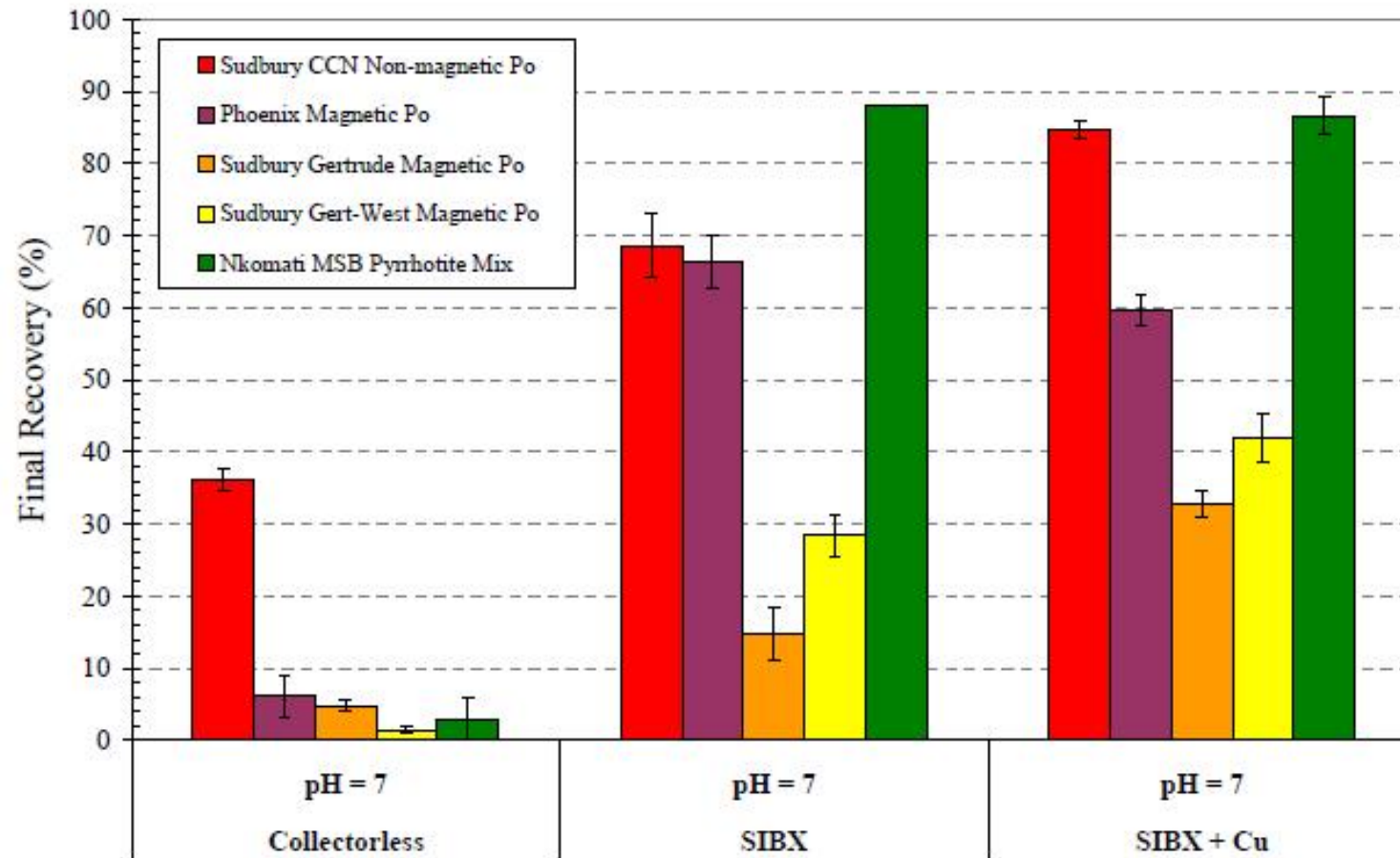
Pyrrhotite Removal - Clarabelle Mill Sudbury



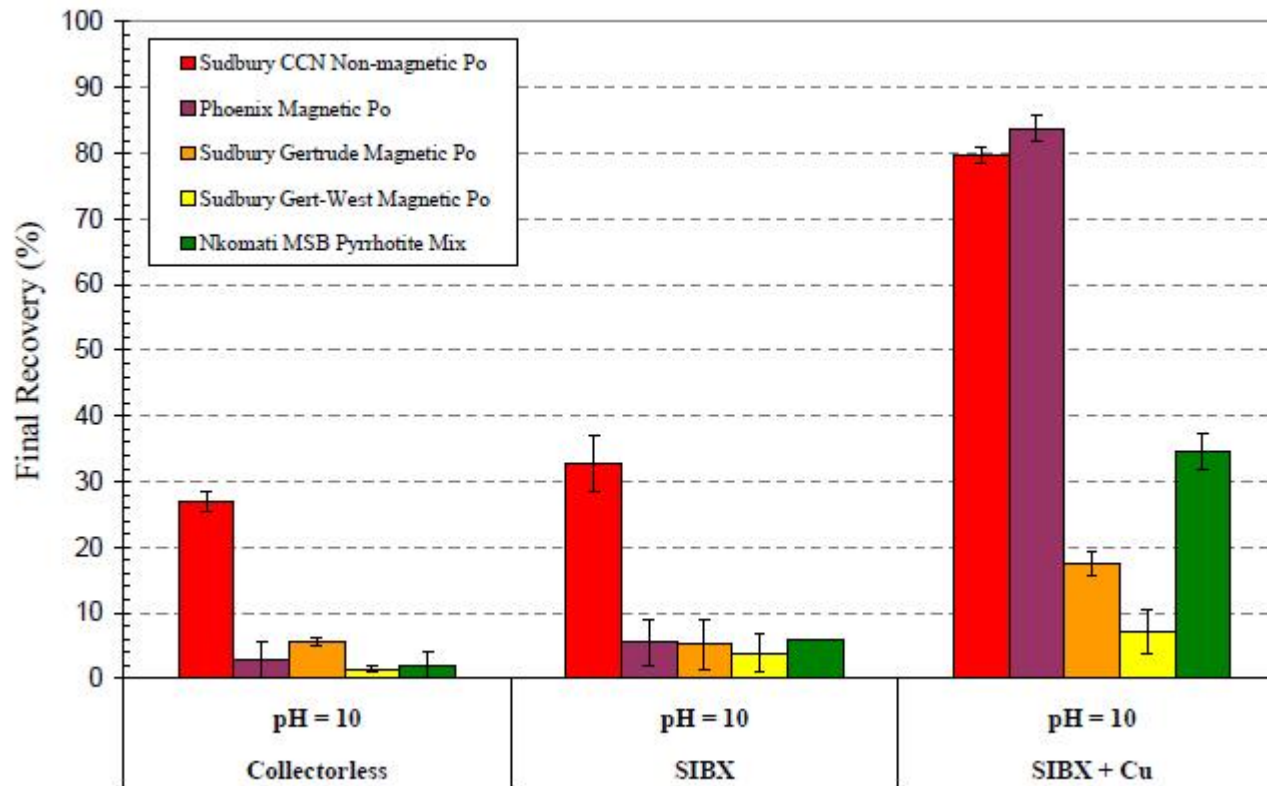
- Magnetic separation can only separate the magnetic pyrrhotite from the pentlandite concentrate

- Flotation is used to remove pyrrhotite from the pentlandite concentrate

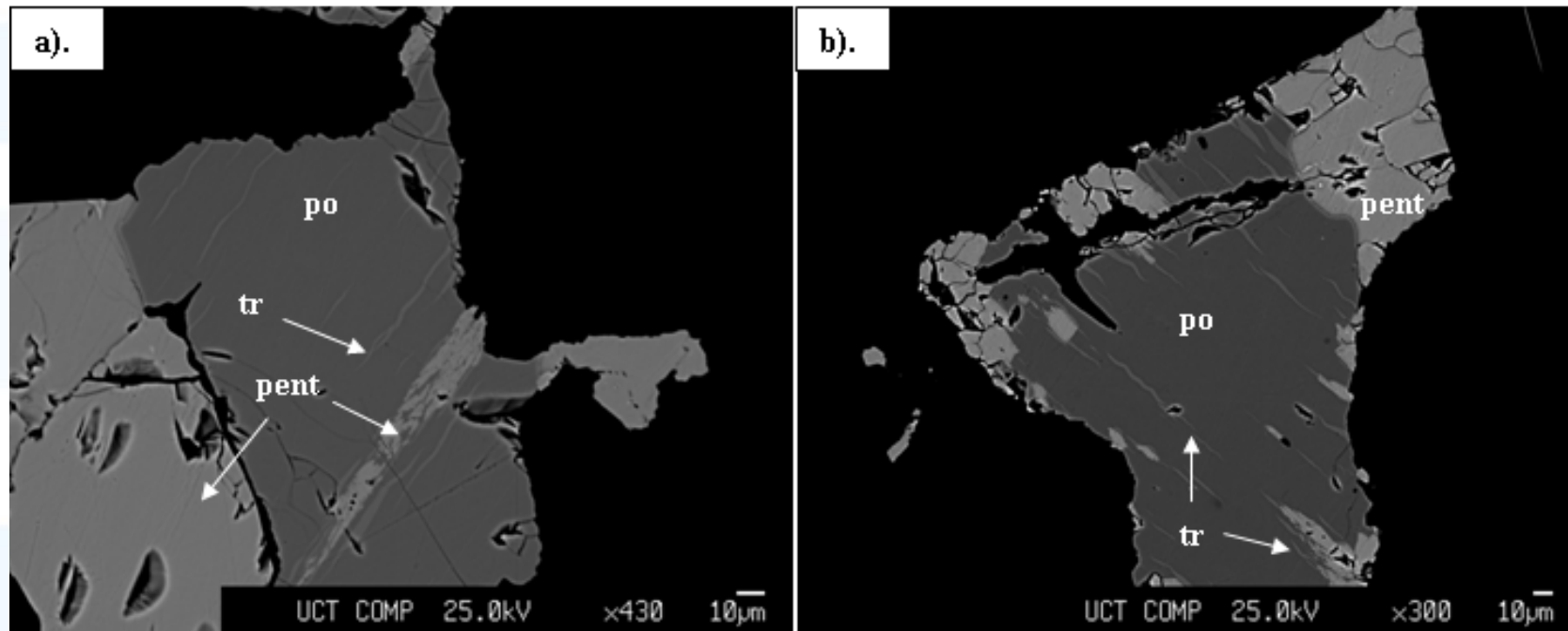
Microflotation Recovery of Pyrrhotite – pH 7



Microflotation Recovery of Pyrrhotite – pH 10



Merensky Pyrrhotite



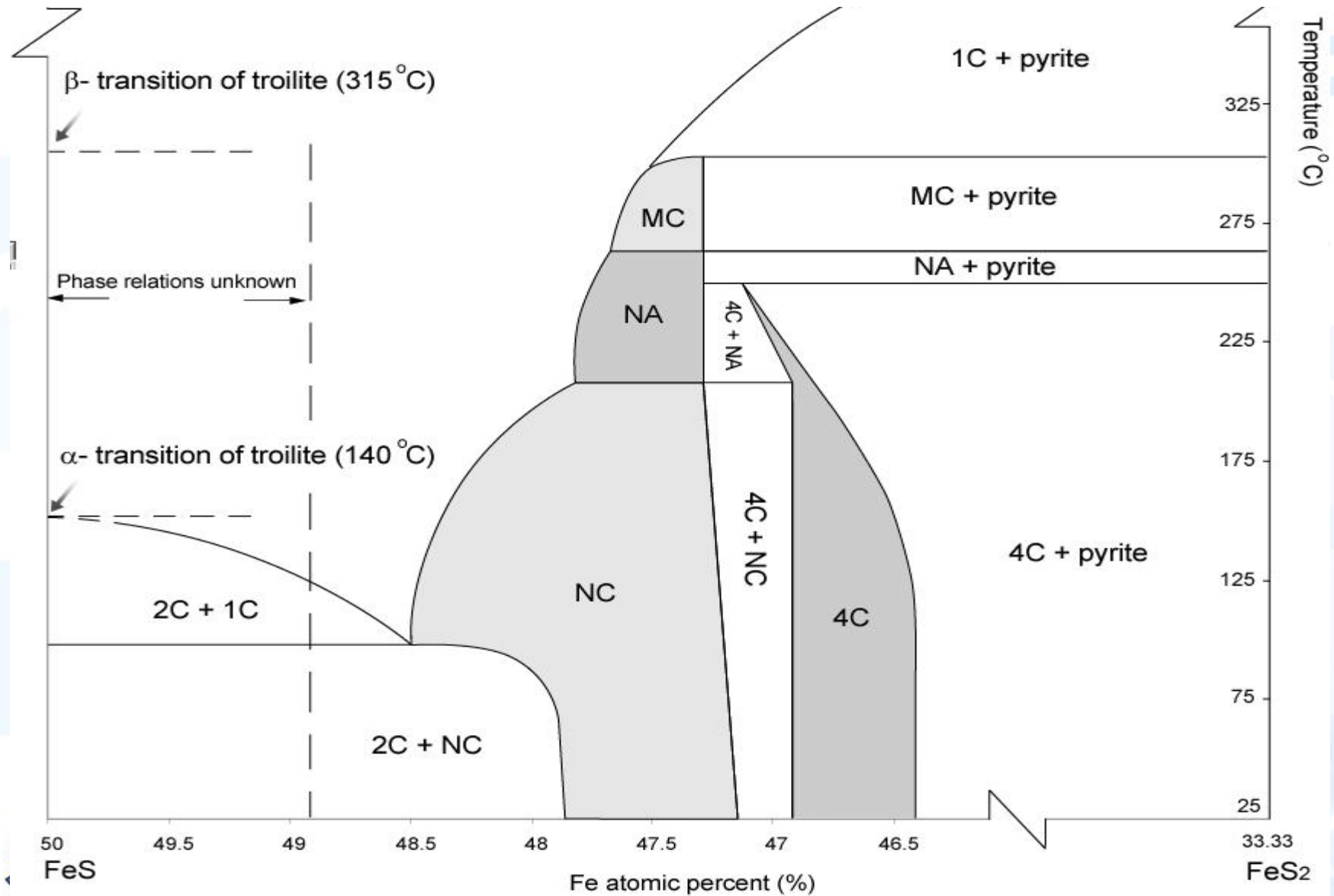
Electron Backscatter Images

Known Pyrrhotites

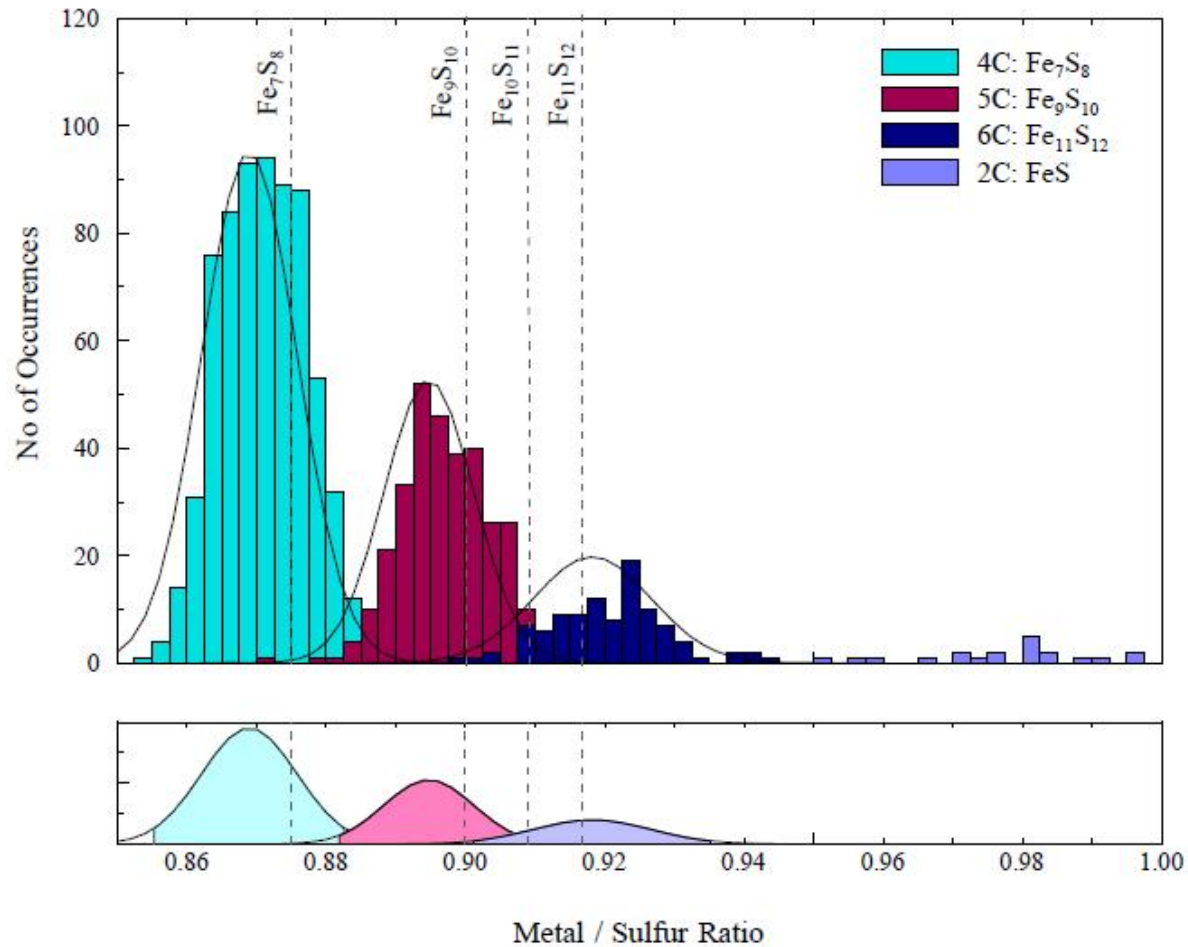
- Troilite FeS (Evans, 1970)
 - Stoichiometric with a 2C superstructure (C is the c-axis of the NiAs subcell)
 - Distorted NiAs structure
- Monoclinic Pyrrhotite (Tokonami et al, 1972)
 - Stoichiometric Fe₇S₈ with 4C superstructure
 - The structure does not explain the non-stoichiometry
- Other natural varieties – structures unknown
 - Fe₉S₁₀ - 5C
 - Fe₁₀S₁₁ – incommensurate
 - Fe₁₁S₁₂ -6C
- Synthetic varieties
 - Fe₇S₈ -Hexagonal 3C
- All varieties thought to be derived from the ordering of vacancies
- Natural varieties have not been synthesized



Pyrrhotite Phase relations



Compositional Variation – 1106 Analyses



New Pyrrhotite structures

- 5C Pyrrhotite from Sudbury, Canada
 - Described as Hexagonal
 - Hexagonal structure determination not successful
 - Structure determined with SHELX and refined to $R = 0.060$
 - Orthorhombic: Cmce
 - $a = 6.893(3) \text{ \AA}$, $b = 11.939(3) \text{ \AA}$ and $c = 28.63(1) \text{ \AA}$
- 6C Pyrrhotite from Mponeng Mine, South Africa
 - Structure correctly predicted by Koto et al. (1975)
 - No atomic coordinates given by them
 - Structure refined with SHELX and refined to $R=0.029$
 - Non-standard Space Group Fd (to preserve orthogonality)
 - $a = 6.897(2) \text{ \AA}$, $b = 11.954(3) \text{ \AA}$, $c = 34.521(7)$, $\beta = 90.003^\circ$
 - Structure can be transformed to Cc space group setting

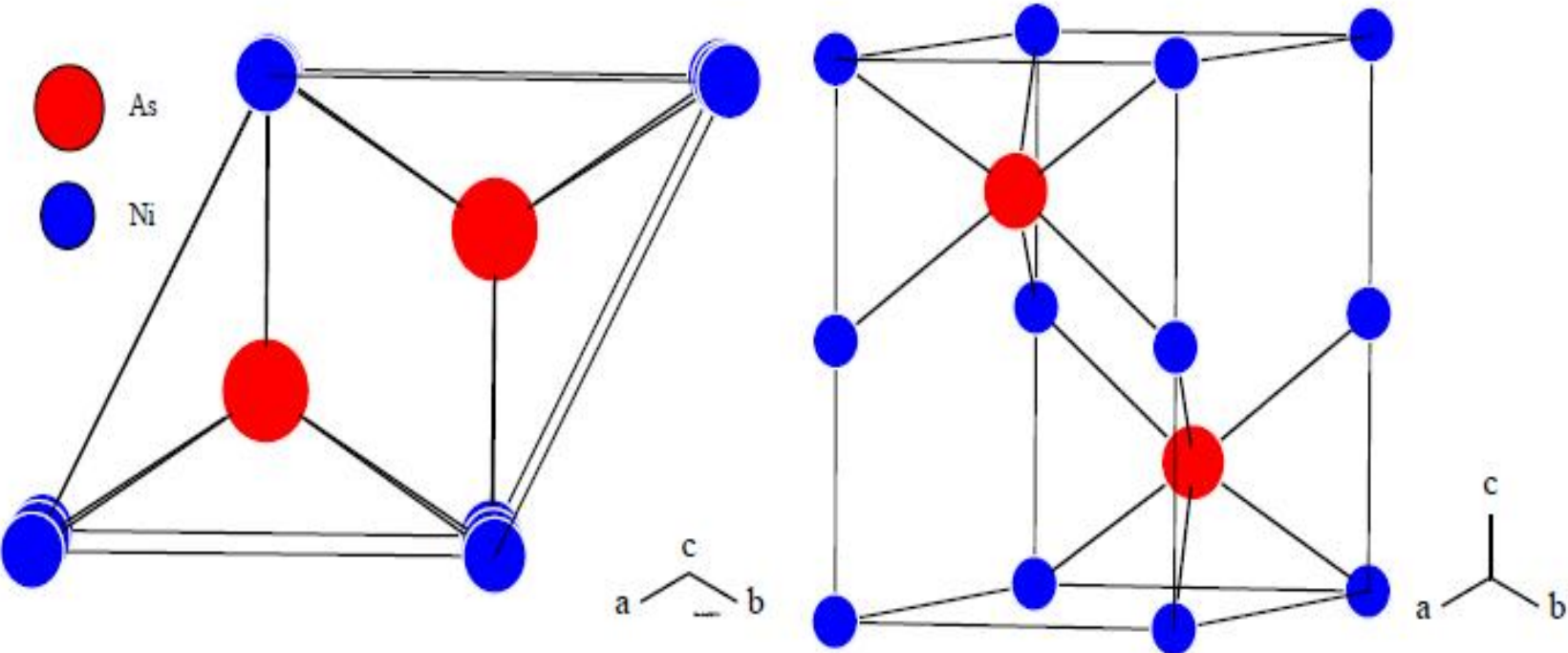


New Pyrrhotite structures

- 4C Pyrrhotite from Bushveld Complex
 - Cell similar to published F-centered cell
 - Cell is C-centered
 - Structure determined with SHELX and refined to $R = 0.052$
 - Monoclinic: C2
 - Cell: $a = 11.890(4) \text{ \AA}$, $b = 6.872(2) \text{ \AA}$, $c = 22.786(8) \text{ \AA}$
 - F2/d: $a = 11.902(8) \text{ \AA}$, $b = 6.859(5) \text{ \AA}$, $c = 22.787(10) \text{ \AA}$
(Tokonami et al.)
- The crystal structures are needed:
 - to understand the compositional variation
 - to quantify the different pyrrhotites
 - to calculate their bonding and oxidation properties
 - to model the docking of reagents on mineral surfaces

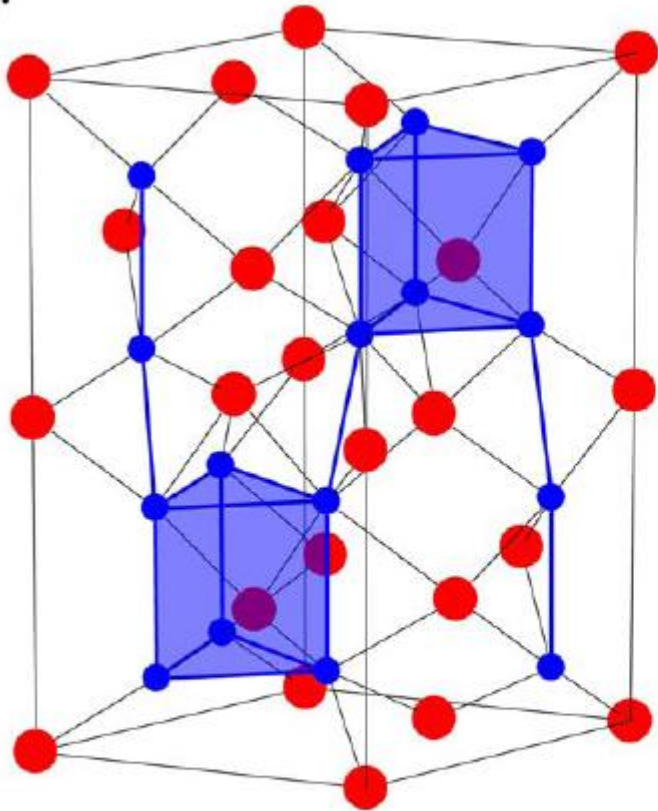


NiAs Structure



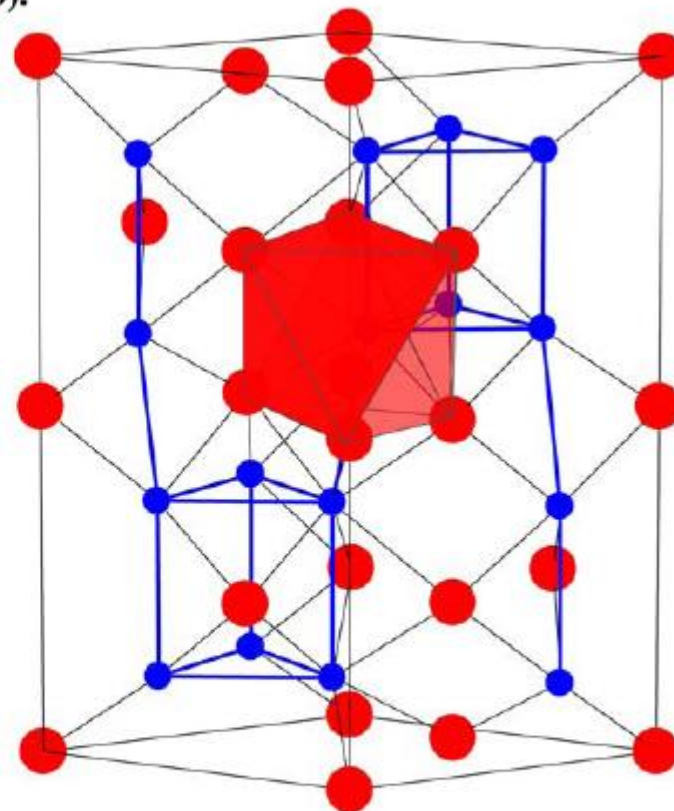
Troilite – FeS (stoichiometric)

a).



a c b

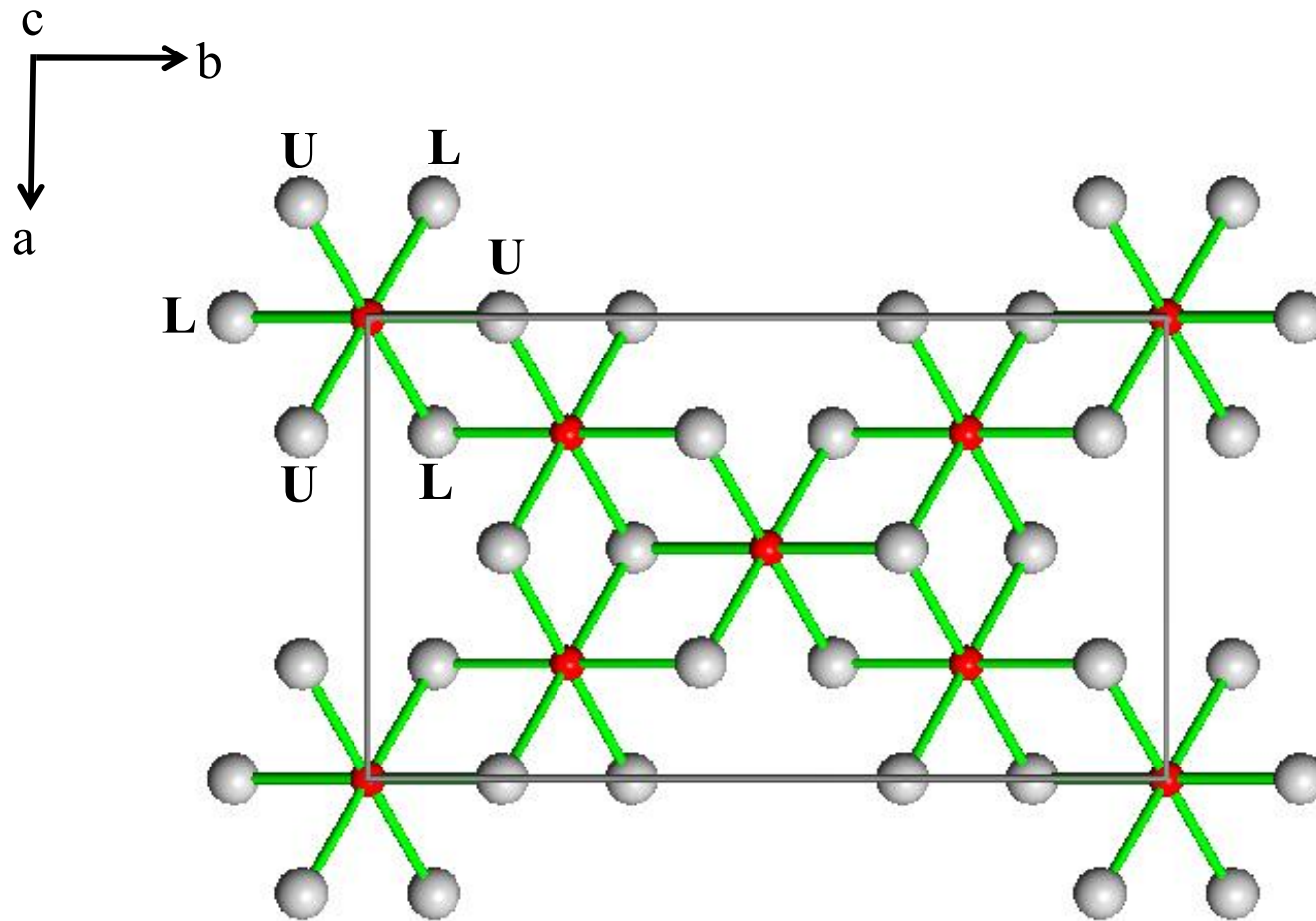
b).



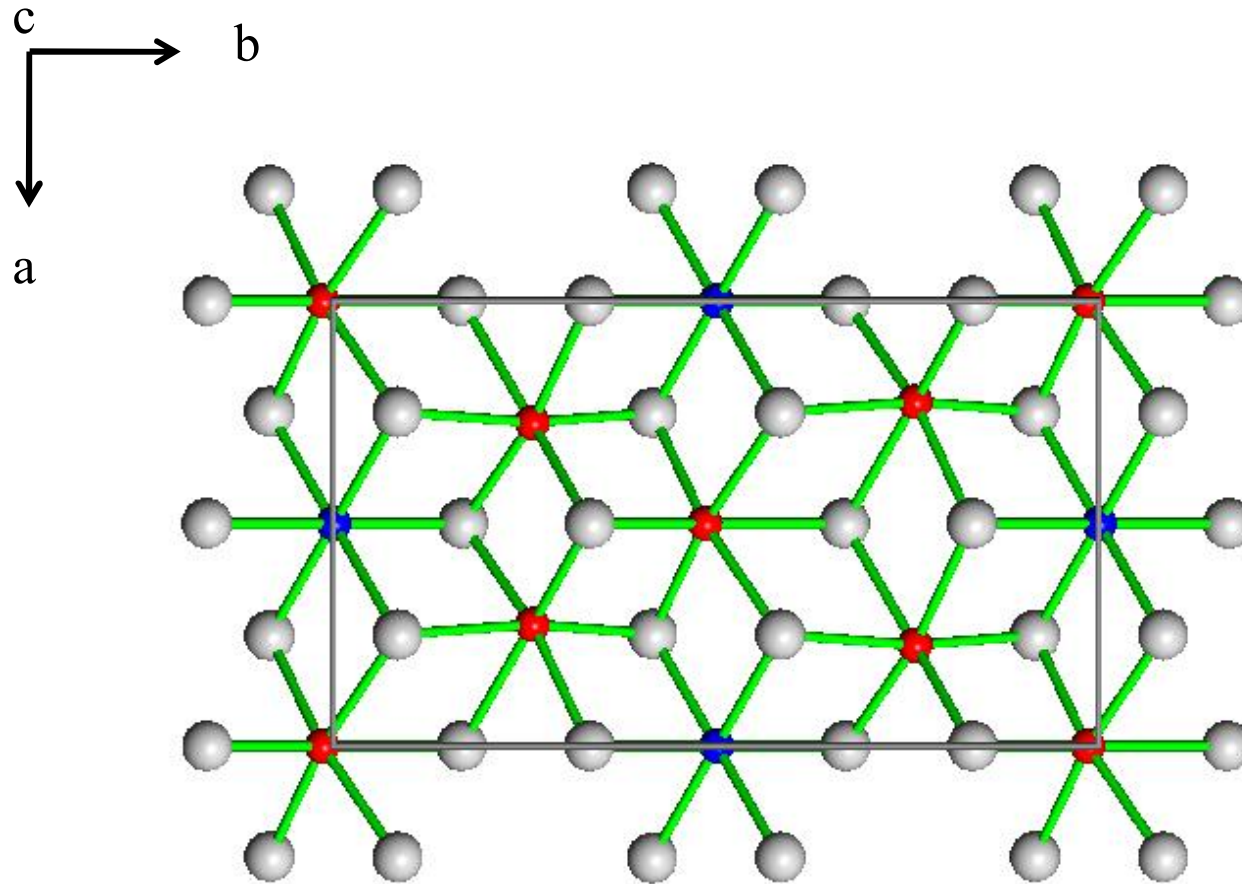
S
Fe



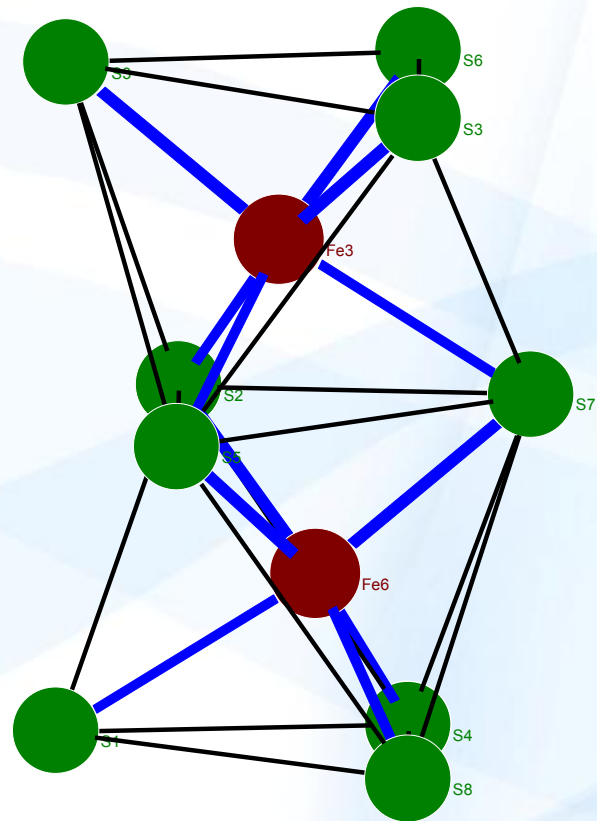
5C Pyrrhotite - Layer-0



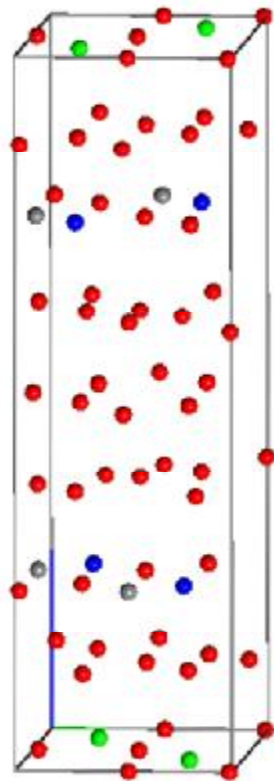
5C Pyrrhotite - Layer-1



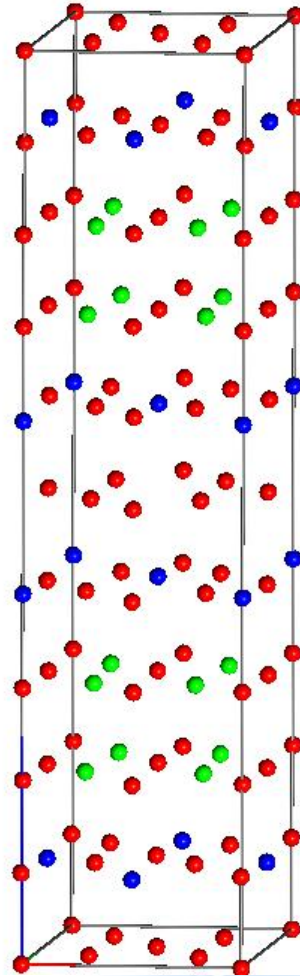
Adjacent Fe- octahedra



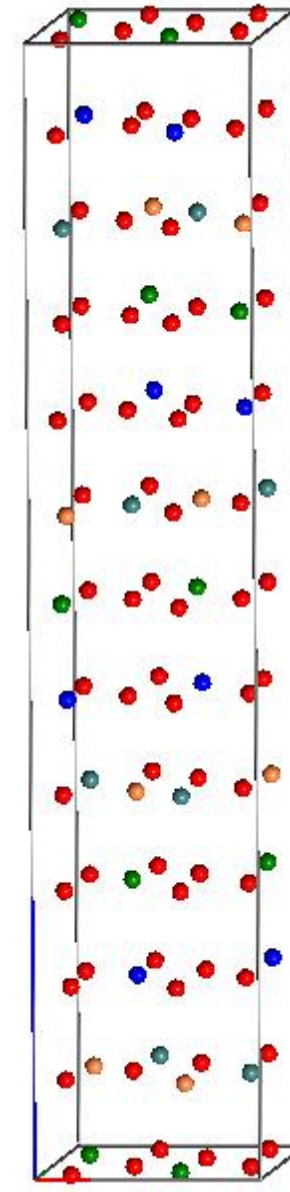
New Pyrrhotite Superstructures



4C



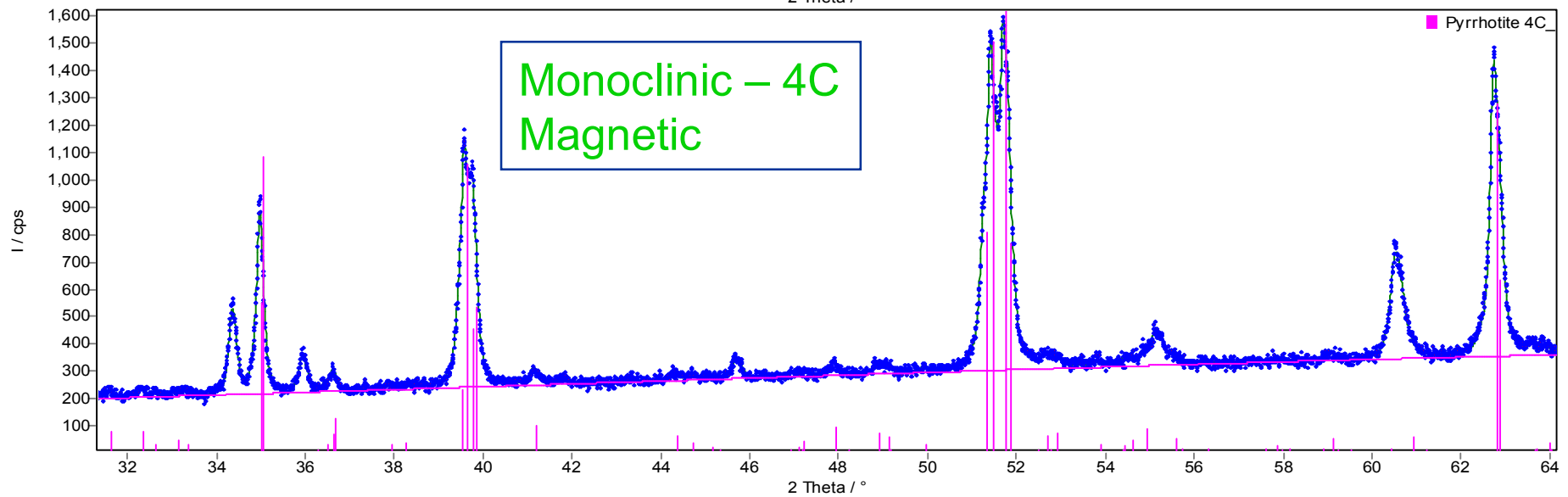
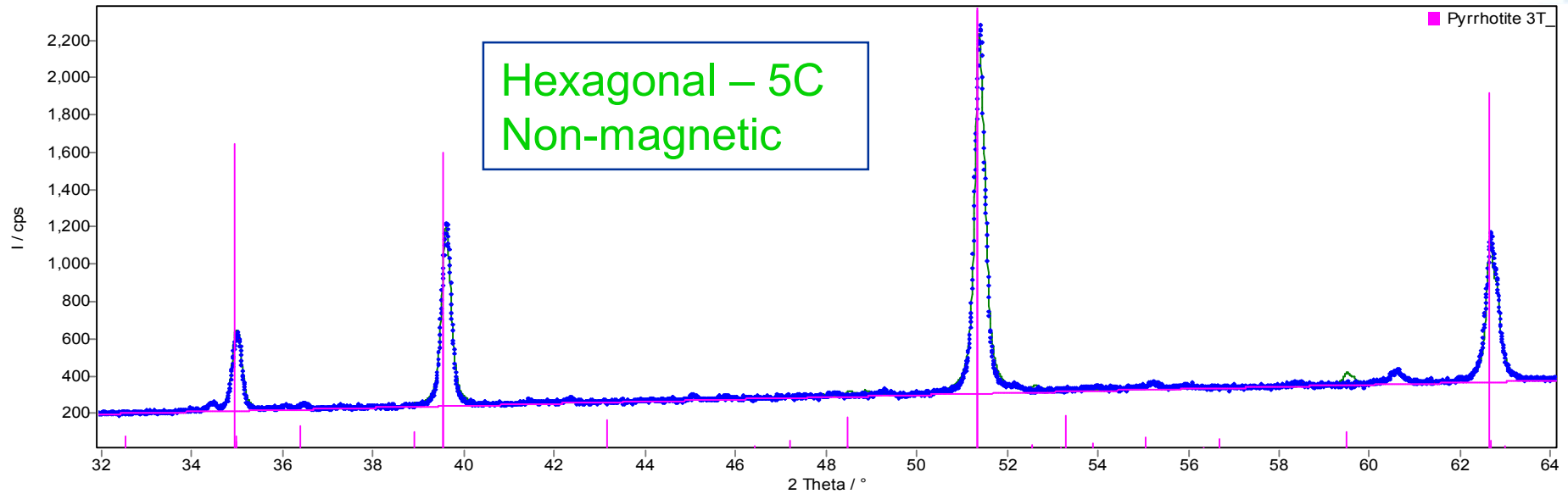
5C



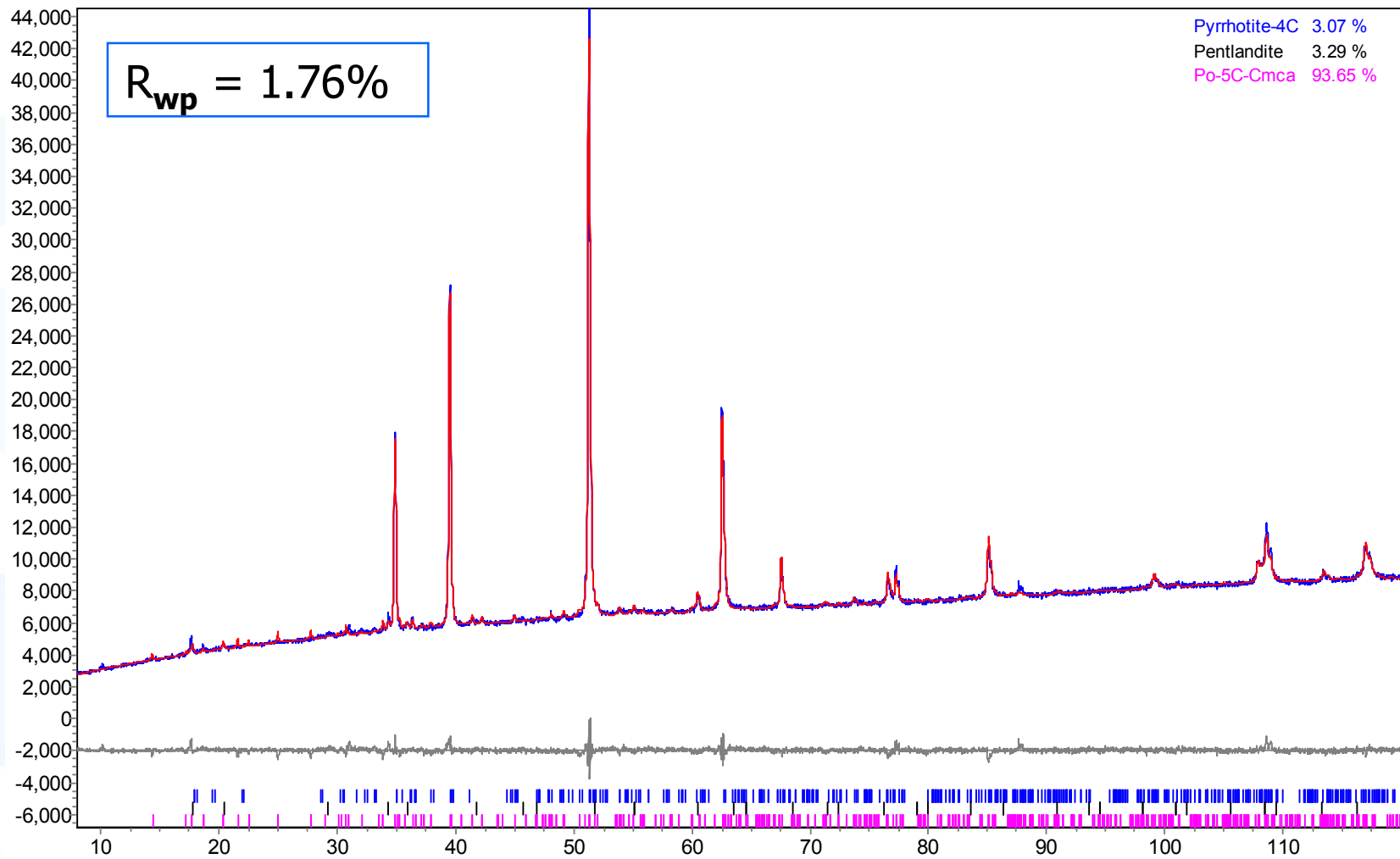
6C



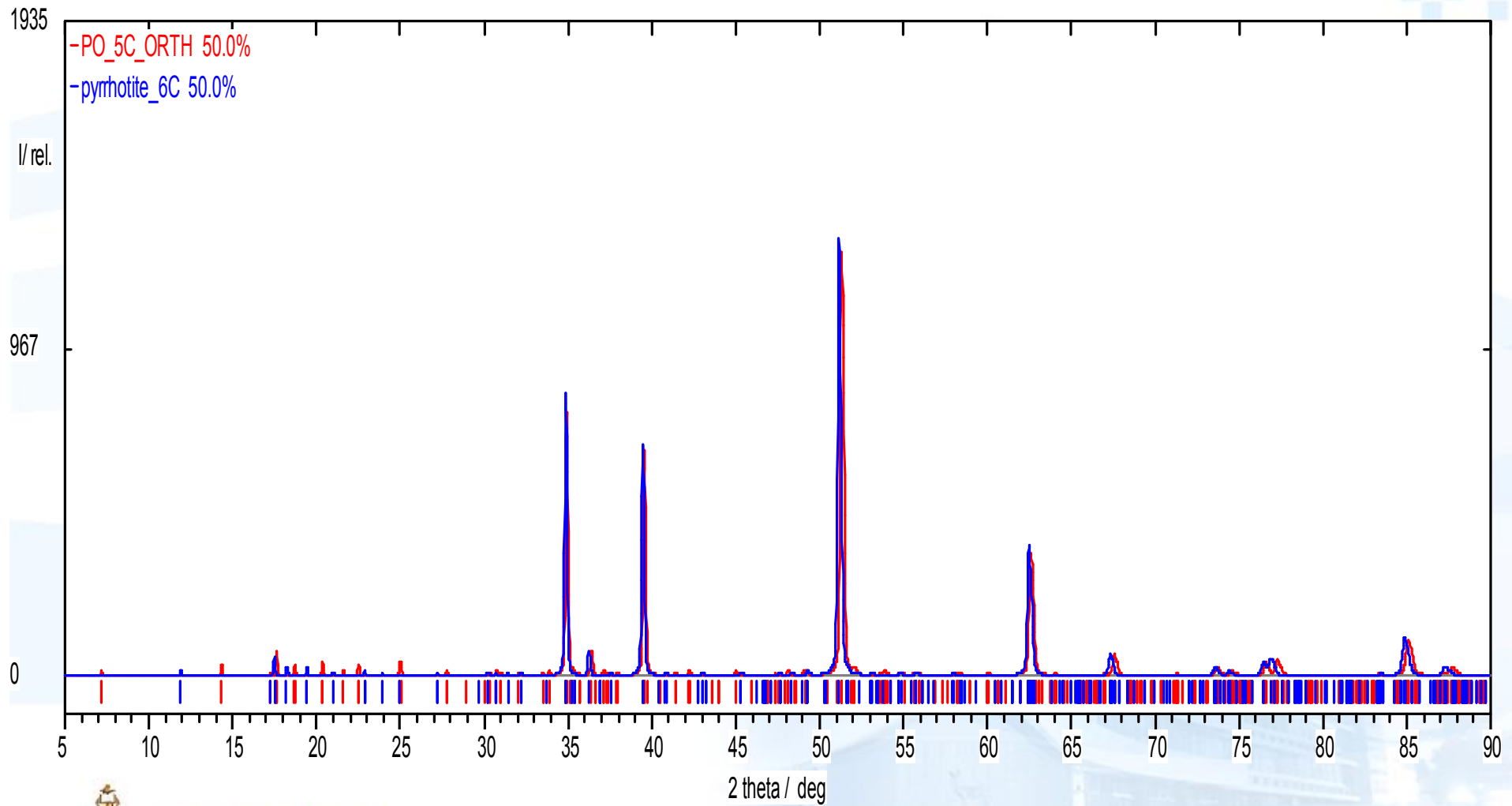
Pyrrhotite - Quantification



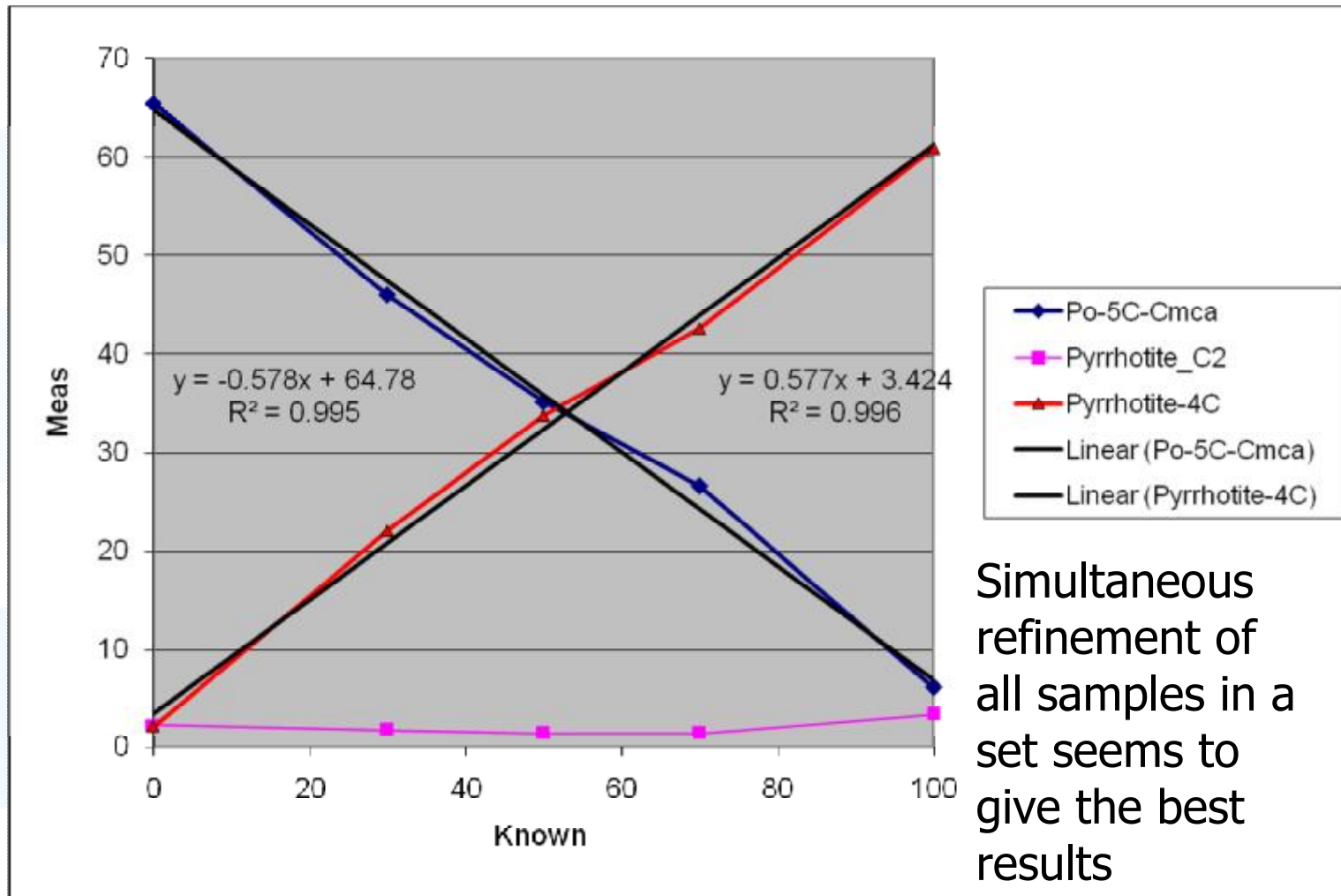
Refinement of 5C pyrrhotite - TOPAS



Calculated XRD Powder Patterns – 5C vs 6C



Accuracy of quantification - Pyrrhotites



Non-stoichiometry

- From the compositional data it seems likely that Fe^{3+} is substituting for Fe^{2+} , i.e. $\text{Fe}^{2+}_5\text{Fe}^{3+}_2\text{S}_8$
- The presence of ferric iron seems to stabilise non-stoichiometric pyrrhotites
- A neutron diffraction study of Fe_7S_8 supports a fast charge transfer with iron valency of $\text{Fe}^{2.29+}$
- No evidence for Fe^{3+} from Mössbauer spectroscopy
- There are probably many more pyrrhotite structures
- The proportion of ferric iron will influence the tendency of pyrrhotite to oxidise:
 - Affecting flotation behaviour
 - Affecting the production of acid mine drainage
 - Affecting combustion in flash furnaces
 - Affecting their magnetic properties



Unresolved questions:.....

- The perceived absence of ferric iron in Mössbauer and XMCD studies of pyrrhotites needs to be resolved
- The crystallographic information of the samples needs to be carefully characterised
- Materials modelling on the structural stability of the superstructures is in progress at UP and Curtin University
- Spectroscopic and neutron diffraction work needs to be done on natural pyrrhotites to resolve their magnetic character
- Powder XRD is not sufficient to characterise natural pyrrhotites – neither are SEM based methods
- Perhaps high-resolution synchrotron and neutron powder data can help
- Flotation behaviour is still unexplained !!!!!
- There is still a lot of work to be done!



Acknowledgments

- Megan Becker for compositional data, for flotation data, and for samples, also for stimulating my interest in pyrrhotites
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